

Dynamics II: Wind-driven Circulation, Cyclostrophic Wind

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1. Ekman layer: Consider a geostrophic flow (\bar{u}, \bar{v})

$$-f\bar{v} = -\frac{1}{\rho_0} \frac{\partial \bar{p}}{\partial x} \quad (1)$$

$$f\bar{u} = -\frac{1}{\rho_0} \frac{\partial \bar{p}}{\partial y} \quad (2)$$

with pressure $\bar{p}(x, y, t)$. The boundary-layer equations are then

$$-f(v - \bar{v}) = \nu \frac{\partial^2 u}{\partial z^2} \quad (3)$$

$$f(u - \bar{u}) = \nu \frac{\partial^2 v}{\partial z^2} \quad (4)$$

The boundary conditions are specified to be at the surface

$$\rho_0 \nu \frac{\partial u}{\partial z} = \tau^x \quad (5)$$

$$\rho_0 \nu \frac{\partial v}{\partial z} = \tau^y \quad (6)$$

and for $z \rightarrow -\infty$: $u = \bar{u}$, $v = \bar{v}$.

a) Calculate the flow (u, v) as the departure from the interior flow (\bar{u}, \bar{v}) !

b) Calculate the net wind-driven horizontal transport through integration

$$V = \int_{-\infty}^0 dz (v - \bar{v}) \quad (7)$$

$$U = \int_{-\infty}^0 dz (u - \bar{u}) \quad (8)$$

What is the direction of U and V in terms of the surface wind stress τ ?

c) For the case $f = f_0$ of constant Coriolis parameter, determine the divergence of the flow

$$\int_{-\infty}^0 dz \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) \quad (9)$$

which is identical to the vertical velocity across the Ekman layer (since $w(0)=0$).

2. Cyclostrophic wind: When the flow is sufficiently near the equator so that f is small or when the Coriolis force is negligible compared to the centripetal acceleration, the gradient wind equation becomes

$$\frac{v \mathbf{k} \times \mathbf{v}}{R} = -\frac{1}{\rho} \nabla_z p \quad (10)$$

where \mathbf{k} is the unit vector in z direction, \mathbf{v} is the velocity vector, v is the meridional velocity, R Earth radius, ∇_z horizontal nabla operator.

a) Derive this equation!

b) What is the associated gradient wind equation including the Coriolis force?

c) What is the Rossby number?